Identification and modeling of nonstationary tendencies in extreme rainfall

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Content

- Data
- Model framework and previous studies
- Temporal development of the extreme rainfall
 - The annual no. of extreme events
 - The α and κ in the Generalized Pareto Distribution
- Conclusion



31 years of observed rainfall from high resolution rain gauges







A regional model of extreme rainfall

- Spatial variations of GP-parameters
- Assumes stationarity
- Rainfall duration 1 min 24 hours
- Model applied in urban drainage design today
- An incorrectly dimension of the drainage system is critical





Temporal variation of λ



Modelling variations of λ

$$\frac{N_{ij}}{t_{ij}} = \phi + \varphi \cdot year_i + \tau \cdot NAO_i + \gamma \cdot MAP_j + \upsilon \cdot latitude_j + \varepsilon_{ij}$$

i=1...31 (temporal variable, number of years)

j=1...70 (spatial variable, number of stations)

N= number of extreme events (a count)

t= active measurement period

 $\lambda = N/t$

Be aware of:

- Correlations between rain gauge stations
- Correlations between the explanatory variables



Variations in λ – sensitivity analysis

Variations in *α* and *κ*



Variations in *α* and *κ*



Variations in *α* and *κ*





How to get a better understanding of the changes in α and κ ?

$$F_t(x) = F(\theta_t, x)$$

$$F_t(x) = \alpha F_1(\psi, x) + (1 - \alpha)F_2(\theta_t, x)$$

$$F_t(x) = \alpha_t F_1(\psi, x) + (1 - \alpha_t) F_2(\theta, x)$$

 $F_1(x)$ Exponential distribution ? $F_2(x)$ Log-normal distribution ?

Summary and conclusion

- The extreme rainfall is non-stationarity
- When *zO* increase we see a tendency towards:
 - Decreasing slopes for the temporal development of λ
 - Increasing slopes for the temporal development of μ
- Difference for different rainfall durations
 - development of λ most clear for 1–60 min
 - development of μ most clear for 60–720 min
- *κ* becomes more negative
- Outlook: Possibility for comparison with transient climate simulations